

THERMAL REMOTE SENSING OF URBAN CLIMATES IN SOUTH AFRICA THROUGH THE MONO-WINDOW ALGORITHM

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ABSTRACT:

Urban Heat Island (UHI) is among some of the challenges plaguing urban environments. There is increase human population within urban environments especially in the developing world, which is a need to understand the climates for their wellbeing. The use of multispectral satellite remote sensing to investigate the climatic conditions through radiation measurement is applied across the two major South African cities. The thermal remote sensing technique applied for this study is the direct determination of land surface temperatures (LST) using multispectral thermal imagery (ETM+). In addition, meteorological data which included air temperature and relative humidity for the same satellite image dates were used. The LST values obtained showed Johannesburg has many micro heat islands scattered across the metro than in Cape Town. These areas of heat islands corresponded to areas of human settlement and more so the unplanned as opposed to the planned ones. The estimated LST values and observed air temperature values with an R^2 of 0.9. It could be concluded that expansion of urban areas in South Africa has led to increased thermal radiation of land surface in densely populated areas.

1. INTRODUCTION

Urbanisation is a phenomenon of change in landuse type by humans for their well being. One of its key aspects is the reduction in vegetation cover and an increase in concreted or paved areas. Increasing urbanisation especially in the developing countries as South Africa, is continuously challenging for town planners and policy makers to incorporate knowledge about climatic conditions within urban spaces and manage towards minimal impacts on human health (Orimoloye et al., 2018). In which case it would be even more difficult for countries like South Africa where these urban spaces are invaded by unplanned settlements which make use of corrugated sheets to construct houses and have no home-garden spaces (Chikoto, 2009). Therefore, it is but essential that studies be conducted for a spatio-temporal appraisal of some climatic parameters such as surface temperature.

One of the important characteristics of urban environments is impervious paved surfaces and reduced vegetation. This characteristic causes more of the sun's incoming radiant energy to be reflected from urban areas and, less of it is converted to latent energy involved in evaporation or transpiration of moisture (Taha, 1997). The larger volume of asphalt, brick, concrete, and other materials in urban areas give it a higher thermal storage capacity than surrounding vegetated surfaces. The urban canopy also stores large amounts of energy during the day which it releases at night (Oke, 1987). Another factor exacerbating relatively high temperatures within urban areas than its surroundings is the additional heat generated by automobiles, air conditioning and other machinery.

The UHI has been described as one of the major challengers to sustainable livelihood within urban areas. The UHI is the excess temperature near the ground (canopy layer) of the central urban locations as being higher than those of nearby or surrounding areas (Voogt and Oke, 2003). This phenomenon has previously been considered as the difference between urban and rural areas but with recent developments in cities, there is no distinct borderline between 'urban' and 'rural' areas as a result of urban

growth (Voogt, 2004). In recent times it has been considered as the difference between the central parts of the city and its surrounding areas (Ngie et al., 2014). The continuous urbanisation of urban areas has led to various settlement areas across metropolitan cities and is a motivation for this study.

Therefore, this study sought to assess the spatio-temporal display of surface temperature across two major metropolitan cities of South Africa with both planned and unplanned settlements. This made use of the mono-window algorithm in estimating LST (Qin et al., 2001) for Johannesburg and Cape Town metros. There will be a highlight of surface temperature situation across these metropolitan cities.

2. STUDY AREA AND DATA COLLECTION

Urban areas in post-apartheid South Africa where movement of people is no longer restricted have experienced tremendous growth rates (Ndebele, 2013). According to the last national census, the country experienced a 15.5 % population growth from 2001 to 2011 meanwhile some provinces like Gauteng (seat to the city of Johannesburg) experienced a 33.7 % increase within the same 10-year period being double the national average (Stats SA, 2012). Based on their rapid rate of urbanisation the two major cities were selected for this investigation (Figure 1).

2.1 Study Area

In terms of population and economic status, Johannesburg is the biggest city in South Africa with huge urban sprawl problems influenced by highly linked highway interchanges. On a land surface area of 1 645 km² the city supports a population of 4 434 827 inhabitants. Cape Town is the second largest with 3 740 026 inhabitants on a total land area of 2 455 km² (census 2011: Stats SA, 2012).

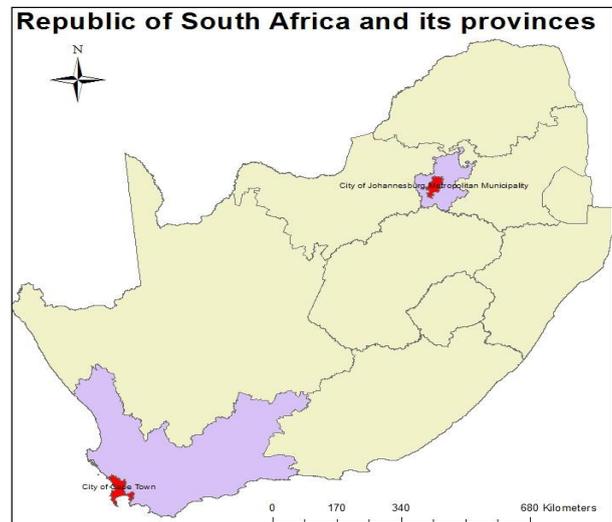


Figure 1: Map of South Africa showing the 2 cities

2.2 Data

2.2.1 Satellite Imagery: Satellite data were acquired over two different seasons: winter (June-August) and summer (January-March) during the day periods for each season for two different years. The years selected were 2002 and 2012. Where suitable 2012 images were not available (due to high cloud cover), 2011 images within the same season were acquired. For instance, summer images over Cape Town were selected in 2011. Therefore, a total of 8 LANDSAT ETM+ images were acquired for the study by downloaded from the United States Geological Society (USGS) website.

2.2.2 Other Auxiliary Data: The weather data were obtained from South African Weather Service. These data included air temperature and relative humidity (Table 1). Daily measurements for the required dates of satellite images were obtained. The air temperature and relative humidity data were used for the analysis.

2.3 Data Pre-Processing

The Enhanced Thematic Mapper Plus (ETM+) data was acquired through the USGS website after a thorough search of available and cloud free scenes for both the winter and summer seasons of the years of study. The downloaded data was unzipped then for those that have no gaps, layer stacking and clipping of study areas were done using ENVI version 5.0¹.

module developed in ERDAS Imagine® was applied. This module operates on two functions; to fetch gaps existing on the given scene for the first image, and to match the histogram for the gaps from the scene of the second image.

The separate image bands were combined into a single multispectral image file to serve as input into the algorithm. This is known as layer stacking and also commonly used to combine image derivatives with spectral bands for further analysis. The images were clipped to the desired study areas (municipal boundaries of the cities which are metropolitan).

The Scan Line Corrector (SLC) which compensates for the forward motion of ETM+ on Landsat 7 failed on May 31, 2003 and recovery efforts were unsuccessful. ETM+ now traces a zig-zag pattern along the satellite ground track with striped images of increasing strip width towards the scene edge. An estimated 22 % of any given scene is lost because of the SLC failure and precise locations of the missing scan lines will vary from scene to scene (Landsat 7 Science Data Users Handbook²).

Therefore, the acquired ETM+ images post 2003 had to undergo gap filling using different scenes of the same Landsat ETM+ sensors from different adjacent dates of acquisition. These dates fall into the same season being investigated. The gap filling

3. METHODOLOGY

3.1 LST retrieval methods

The UHI effect has been studied through the retrieval of LST from satellite remote sensing data (Voogt & Oke, 2003; Mirzaei & Haghghat, 2010; Ngie *et al.*, 2014; Abutaleb *et al.*, 2015; Ngie *et al.*, 2016; Sharma & Bisht, 2019). Landsat TM and ETM+ have a single thermal band (band 6) which is the thermal infrared (TIR) channel that records radiation within a 10.4- 12.5 μ m spectral range. Different LST retrieval methods have been developed according to different data sources (thermal bands on a sensor), such as the split-window, temperature/emissivity separation, mono-window, and the single-channel methods (Qin *et al.*, 2001, Sobrino *et al.*, 1996; Jimenez-Munoz & Sobrino, 2003). Among these methods three can be applied to Landsat data; these include the radiative transfer equation, mono-window algorithm and single-channel algorithm.

Although all of these methods can provide good results, some of them such as the radiative transfer equation can only yield results with *in situ* atmospheric parameters captured simultaneously as the satellite passes over the study area. In

addition, the mono-window algorithm can get better results than the single-channel algorithm with a root mean square deviation of 0.9°C (Sobrino, *et al.*, 2004). It has proven to yield better accuracy for the retrieval of LST from Landsat images as illustrated for Hong Kong with a 0.7°C error (Liu & Zhang, 2011).

This study therefore makes use of the mono-window algorithm to retrieve the LST from the ETM+ imagery for the cities of Johannesburg and Cape Town. The algorithm is based on thermal radiance transfer equation and only requires three parameters: emissivity, transmittance and effective mean atmospheric temperature to retrieve LST from Landsat multispectral imagery (Qin *et al.*, 2001). The model was designed using ERDAS Imagine® spatial model tool (Sun *et al.*, 2010; Liu & Zhang, 2011). It required two measured parameters: near-surface air temperature and relative humidity (RH) simultaneously measured as the satellite passed over the study area. These data were obtained from the local weather stations situated within the study areas (Table 1). The two parameters were then converted to atmospheric transmittance and effective mean atmospheric temperature. The third parameter (emissivity) was calculated from the normalized difference vegetation index (NDVI).

¹ Exelis Visual Information Solutions. 2010. Name of the manual. Boulder, Colorado: Exelis Visual Information Solutions.

² <http://landsathandbook.gsfc.nasa.gov/> Accessed 19/11/2016

CITY									
Johannesburg	January 7 th 2002	22							
	July 2 nd 2002	6	5						
	January 3 rd 2012	24	21						
	June 27 th 2012	7	6	50					
Cape Town	March 15 th 2002	26	13	89	97				
	July 21 st 2002	24	13	42	84	1.275	0.8		
	March 26 th 2012	28	15	64	93	2.033	0.811	0.787	
	June 12 th 2011	20	2	74	99	1.214	0.877	0.865	0.891 0.8

Table 1: Satellite and auxiliary data collected and estimated parameters (water vapour content and transmittance)

3.2 Mono-window algorithm

3.2.1 Conversion of digital number into spectral radiance: The satellite sensors measure reflectance from the earth surface as digital numbers (DN) representing every pixel of the image. The DN data from Landsat was converted to spectral radiance using Equation (1):

$$L_{\lambda} = L_{\lambda min} + (L_{\lambda max} - L_{\lambda min}) \times \frac{(Q_{\lambda DN} - Q_{\lambda min})}{(Q_{\lambda max} - Q_{\lambda min})} \quad (1)$$

Where L is at-sensor spectral radiance; * $L_{\lambda max}$

L is maximum at-sensor spectral radiance; $L_{\lambda min}$

L is minimum at-sensor spectral radiance;

($Q_{\lambda max}$ Q= 255) and ($Q_{\lambda min}$ Q= 1) are the maximum and minimum DN value of pixels respectively and, $Q_{\lambda DN}$ Q is the DN value of pixel.

For Landsat ETM⁺ data, equation (1) can be rewritten in two dimensions depending on the low and high gain bands of the thermal band and expressed in equations 2 and 3.

- 1 The low gain band which provides an expanded dynamic range with lower radiometric resolution (sensitivity), has less saturation at high DN values³ and expressed thus:

$$L_{61} = 0.0671(Q_{\lambda DN} - 1) \quad (2)$$

- 2 The high gain band has a higher radiometric resolution (sensitivity) although it has a more restricted dynamic range (*ibid*) and expressed thus:

$$L_{62} = 0.0372(Q_{\lambda DN} - 1) + 3.2 \quad (3)$$

3.2.2 Conversion of spectral radiance into at-sensor brightness temperature: The conversion formula is:

$$T_i = \frac{K_2}{\ln\left(\frac{K_1}{L_{\lambda}} + 1\right)} \quad (4)$$

³http://landsathandbook.gsfc.nasa.gov/data_properties/prog Sect6_4.htm. Accessed 27/02/2014

Where T is at-sensor temperature in Kelvin, K_1 and K_2 are pre-launch calibration constants and, L_{λ} is the spectral radiance in watts per metre squared steradian micron ($W/(m^2 \cdot sr \cdot \mu m)$). For ETM⁺ the calibration constants are:

$$K_1 = 666.09 (W/(m^2 \cdot sr \cdot \mu m))$$

$$K_2 = 1282.71 K$$

3.2.3 Calculation of NDVI: NDVI is a simple graphical indicator usually used to analyze remote sensing measurements, and assess if observed target includes green vegetation or not. NDVI is one of the outputs that run into the model to retrieve LST. Therefore, eight NDVI maps were produced according to the various dates of satellite capture in this study using equation (5)

$$NDVI = \frac{NIR - R}{NIR + R} \quad (5)$$

Where NIR is Near Infra-red band and R is the red band

3.2.4 Estimation of emissivity: In this study, the land surface emissivity was derived from the NDVI (Zhang, 2006). From the NDVI results obtained above (Equation 5) corresponding land emissivity values were established (Table 2).

NDVI	Land surface emissivity (ϵ_i)
NDVI < -0.185	0.995
-0.185 ≤ NDVI < 0.157	0.970
0.157 ≤ NDVI ≤	1.0094+0.004

Table 2: Estimation of land surface emissivity from NDVI (Adopted from Zhang, 2006)

3.2.5 Estimation of atmospheric transmittance: Atmospheric transmittance which is defined as the capacity of the atmosphere to transmit electromagnetic energy is dependent on the air mass penetrated by radiations, as well as on the amount of water vapour and dust in the air. It varies for radiations of different wavelengths; the smaller the atmospheric absorption and scattering of light, the greater the atmospheric transmittance.

3.2.6 Calculation of water vapour content: The calculation of water vapour is done through equation (6) which entails deriving relative humidity. Relative humidity can be defined as the ratio of water vapour at a given temperature to the saturation water vapour pressure at the same temperature (Wagner and Pruß, 2002).

$$RH = \frac{P_a}{P_s} \times 100 \quad (6)$$

Where RH is the relative humidity;

P_a is the actual water vapour pressure;

P_s is the saturation water vapour pressure.

Equation (6) is rewritten to obtain P_a as follows:

$$P_a = P_s \frac{RH}{100} \quad (7)$$

The saturation water vapour pressure (P_s at a given temperature (K) can be calculated (Wagner and Prueß, 2002):

$$P_s = P_c * e^{\frac{T_c}{T} (C_1 \theta + C_2 \theta^{1.5} + C_3 \theta^3 + C_4 \theta^{3.5} + C_5 \theta^4 + C_6 \theta^{7.5})} \quad (8)$$

Where:

T = ambient temperature in K

P_s = Saturation water vapour pressure (hPa)

T_c = Water critical temperature (647.096 K)

P_c = Critical pressure (220 640 hPa)

C_1 = -7.85951783

C_2 = 1.84408259

C_3 = -11.7866497

C_4 = 22.6807411

C_5 = -15.9618719

C_6 = 1.80122502

And

$$\theta = 1 - \frac{T}{T_c}$$

Atmospheric transmittance is then estimated on the basis of some linear equations derived from statistical regression analysis (Qin *et al.*, 2001). These equations relate the atmospheric transmittance to the water vapour content (g/cm^2) (Equation (8)). Based on the results obtained for water vapour content in this study, corresponding atmospheric transmittance equations were selected (**Error! Reference source not found.**). The air temperature (t_a) profiles were set as $t_a < 15^\circ \text{C}$ being low and $t_a > 15^\circ \text{C}$ as being high.

3.2.7 Calculation of mean atmospheric temperature:

Equations (9) and (10) are used to estimate the atmospheric mean temperature (T_a) based on given near surface air temperature (Qin *et al.*, 2001):

For mid-latitude summer

$$T_a = 16.0110 + 0.92621T_0 \quad (9)$$

For mid-latitude winter

$$T_a = 19.2704 + 0.91118T_0 \quad (10)$$

LST retrieval from the mono-window algorithm: In the mono-window algorithm, the LST value is calculated using Equation 11 but with two unknowns (C and D) being realised through Equations 12 and 13. The three required parameters (emissivity, transmittance and effective mean atmospheric temperature) are included in the equation as follows:

$$T_s = \frac{(a+bT_i)(1-c-D)+D(T_i-T_a)}{c} + T_i \quad (11)$$

$$D = (1 - \epsilon_i)[1 + (1 - \epsilon_i)t] \quad (12)$$

$$C = \epsilon_i * t \quad (13)$$

Where: T_s is LST (K); T_i is at-sensor temperature (K); ϵ_i is emissivity estimated from NDVI; t is atmospheric transmittance calculated from water vapour content; T_a is mean atmospheric temperature; $a = -67.355351$, and $b = 0.458606$.

4. RESULTS AND DISCUSSION

4.1 Relationship between the observed air temperature and the estimated LST

Statistical regression analysis was used to establish a relationship between the estimated LST values and the measured air temperature values from weather stations in the two cities. The analysis was done using mean temperature values for both scenarios over the cities and seasons. The results yielded an R^2 of 0.9. Therefore, results confirmed the fitness of the mono-window algorithm in estimating LST values. The difference between the estimated LST values and the measured air temperature values can be associated partly with the effects of surface roughness on surface temperature and emissivity not considered (Weng, 2001) and partly because of atmospheric impurities which can obstruct the smooth passage of radiant energy (Brooks, 2006).

4.2 Spatial distribution of surface temperature

The spatial distribution of surface temperatures from these two cities indicated that not just do the central city regions illustrate heat islands but also around the cities which correlated to the informal human settlement areas. The intensity of these heat islands varied within these cities and the various areas.

In Cape Town the city centre showed a surface temperature increase to a maximum of 5°C compared to the surroundings in summer 2012 (Figure 2). However, the geographical features surrounding this city being the Table Mountain in the south and the ocean (Atlantic and Indian Oceans) to the east and west respectively should play a role in the experience but require further research to ascertain.

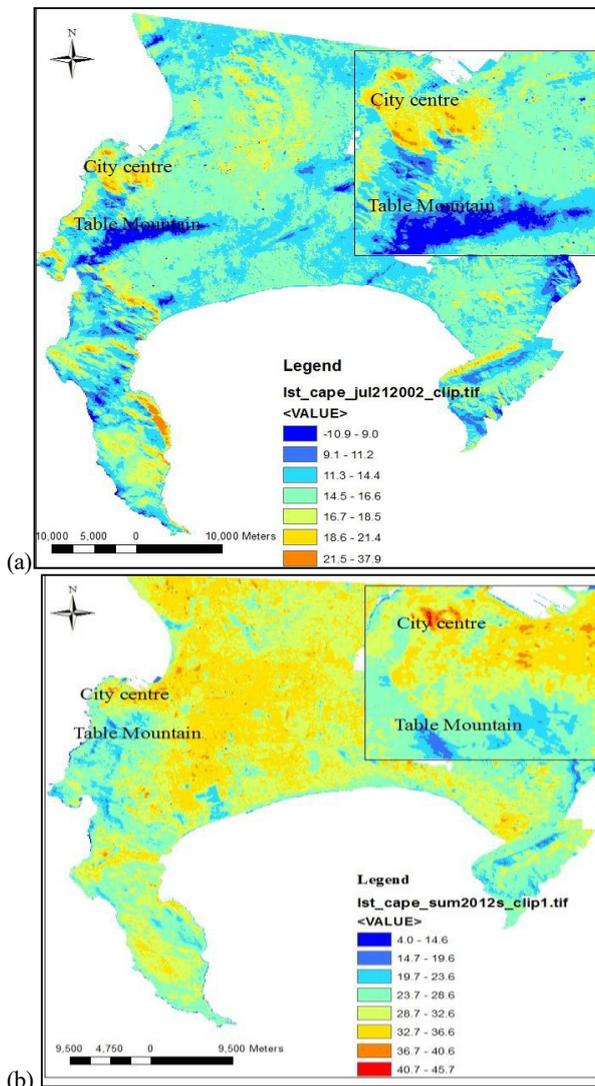


Figure 2: LST maps for Cape Town (a) Winter 2002; (b) Summer 2012

When new surface materials which include asphalt and concrete are introduced to a naturally vegetated area the thermal climate also changes with an increase in canopy temperature. The emission of heat, moisture and pollutants dramatically change radiative, thermal, moisture, roughness and emissions properties of the surface and the atmosphere above (Joshi & Bhatt, 2012). There is also the emission of a lot of radiant energy through automobiles and heavy machines within these urban centres which contributes in the higher LST values (Figure 3).

Generally, the intensity of surface temperature in Johannesburg is at micro levels due to the high density of planted trees which do not only absorb the energy but cause shaded areas that disrupt the heat island (Liu *et al.*, 2007). There are other major centres of heat islands in Johannesburg which could be linked to the expansion of urban areas with or without necessary planning. The expansion of planned areas towards the northern parts of the city, distinctively shows the Sandton, Fourways and Randburg city centres. However, the unplanned and unorganized rapid expansions of settlements such as the Alexandra Township in the north as well and SOWETO in the south western section illustrate heat island intensities slightly higher than their planned human settlements (Figure 3). For instance, Sandton which is adjacent to Alexandra but is a

formally planned up-market area experiences a 2° C lower land surface temperature.

This study highlighted the fact that it is not just a challenge with development that has transformed vegetated areas to paved or concreted land covers but also that of land use. The human activities within these unplanned settlements which include burning coal for cooking as well as heating in open fires can be seen as contributors to raising the surface temperature. These unplanned settlements have limited land space for vegetation due to the high influx of people from rural to urban areas without any resources to live within decent or planned neighbourhoods (Chikoto, 2009). They end up raising structures with metal sheets and of high density as is the case with the informal sections of the south western part of Johannesburg (SOWETO), Alexandra and other associated townships in the north eastern sections contributing to the higher LST values over those areas (Figure 3).

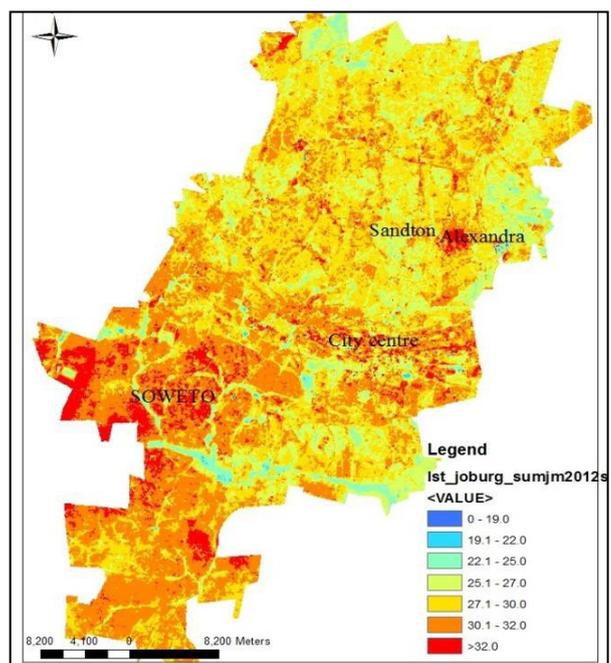


Figure 3: LST map of summer 2012 for Johannesburg

Another interesting observation for the city of Johannesburg is that the presence of mine tailings offer a cooling effect within the city centre as illustrated by the LST values over them (Figure 3). However, the rehabilitation process (introduction of vegetation) of these tailings has provided more assistance in regulating the urban temperatures. Li *et al.* (2011) studied the relationship of LST to NDVI and vegetation fraction and indicated that subsurface temperatures are linked to land-use and the vegetated areas act as heat sinks.

Human settlements have replaced existing vegetation with paved surfaces, which results in the sun's energy heating the man-made surfaces rather than being absorbed by vegetation and used for evapotranspiration (Taha, 1997) is considered prime cause of the UHI effect over these cities. The hottest temperatures were generally found in areas with the least vegetation (NDVI) and the greatest urban development. There is a strong relationship between LST and impervious surfaces as revealed in this study. This contributes to a large extent of the heat island problem thereby suggesting that town planners need to structure built-up areas with diverse structures and

interspersed vegetation as well as water bodies in varying percentages to serve as the heat sinks and regulating urban climates (Zhang *et al.*, 2008).

There also exist some situations of anomalies in this study where the city centre of Johannesburg seems to be cooler than its surrounding areas in winter but with a well-developed UHI in summer (Figure 3). This is however, not strange as it confirms some empirical generalizations offered by Oke (1982). These generalizations include the UHI intensity being most well developed in summer or in the warm half of the year and that the UHI may disappear by day or the city may be cooler than the rural (surrounding) environs. However, there will be need for further research to investigate the factors contributing to these situations.

5. CONCLUSION

The mono-window algorithm provided a simple and highly effective method for retrieving land surface temperatures for the analysis over these cities. This affirms other similar studies elsewhere in the world which have applied this algorithm (Ifatimehin & Adeyemi, 2008; Lu *et al.*, 2009; Liu & Zhang, 2011; Khaled *et al.*, 2015). The LST derived from the ETM+ thermal band depicted the spatial variability of land surface temperature of the cities of Johannesburg and Cape Town.

The spatial variability and intensity of the surface temperatures across the cities especially in Johannesburg illustrated the contribution from high urbanisation experience. There was a further distinction in the planned and unplanned or poorly planned expansions of urban centres. The intensities did not only vary by space or land use or land cover but also over seasons wherein for the Johannesburg the phenomenon is present in summer and absent in winter. There is a need to further look into the impact of geographical features around these cities such as mountains or oceans and also more validation stations to the high spatial variability which was supported by only two weather stations.

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